Parameterization of momentum flux and energy flux associated with orographically excited internal gravity waves in a baroclinic background flow

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ABSTRACT. A linear dynamical model for airflow across a three-dimensional meso-scale elliptical barrier has been used to parameterize the vertical energy flux ($E_z$) and the horizontal components of momentum flux ($\tau_{zx}$ and $\tau_{zy}$) associated with orographically excited internal gravity waves (IGW) in a baroclinic background flow. This model can be used for an arbitrary orographic barrier, although in the present study an elliptical barrier has been considered. Realistic vertical variations in the stability have been incorporated in the model. The solution of model has been obtained quasi-numerically for some selected cases near the Indian mountains [The Western Ghats (WG) and the Khasi-Jayantia (KJ) hills].

(i) Study shows that the proposed model can compute the energy flux ($E_z$) and the horizontal components of momentum flux ($\tau_{zx}$ and $\tau_{zy}$) associated with orographically excited gravity waves at each level for any barrier.

(ii) Computations for the selected cases shows that the fluxes vary in the vertical. The vertical variation is not uniform with height. It depends on the vertical profiles of the wind and temperature of basic flow for respective cases. However,
the influence of the ‘V’ component (parallel to major ridge axis of elliptical barrier) of basic wind has been observed in all cases.

(iii) For the WG, $\tau_{xz}$ is invariant in the vertical in absence of ‘V’ component. Other fluxes are invariant in the vertical in some layers. Whether this component is present or not, an upper layer is found where all the vertical fluxes are non-divergent.

(iv) For the KJ hills, in absence of ‘V’ component $\tau_{zx}$ and $\tau_{zy}$ both are invariant in the vertical, presence of ‘V’ component makes $\tau_{zx}$ and $\tau_{zy}$ divergent/convergent.

Key words – Orographic gravity wave, Momentum flux.

1. Introduction

We know that, when a stably stratified air-stream flows across an orographic barrier, gravity waves are excited which propagate upwards under certain conditions of thermal stability and airflow stratification. These orographic gravity waves can transport momentum from a stably stratified airstreams to the earth’s surface in response to a net pressure drop between the windward and lee slopes of the barrier. It is also known that these gravity waves are also capable of transporting energy from surface to the mean flow at great height. It is also believed that the continuous extraction of momentum from the mean flow is one of the possible causes for breaking of gravity waves. But this transport of momentum or energy by orographic gravity waves is a sub-grid scale phenomenon, yet its impact on large-scale motion is substantial. So the parameterization of these sub-grid scale fluxes is important in the numerical weather prediction models.

Sawyar (1959) first pointed out the relative importance of this momentum loss of the mean flow due to continuous extraction of momentum from it by the orographic gravity waves. He examined the case of a two-dimensional (2-D) flow over a bell-shaped obstacle with half width (a) 2 km and height (b) 300 m and determined that the typical value of wave momentum flux is of the order 1-10 dynes/cm².

Eliassen and Palm (1961) showed that for 2-D linear gravity waves, the vertical flux of horizontal momentum, resulting from waves, is independent of height when the waves are steady and non-dissipative.

Bluemen (1965) noted that the magnitude of the wave drag is sensible to the vertical wavelength. He also showed that the maximum value of the drag is attained when the vertical wavelength is twice the maximum height of the mountain.

Brettherton (1969) reviewed the theories concerning the propagation of internal gravity waves (IGW) in a horizontally uniform shear flow. His computations showed that for a 19m/s gradient wind over hilly terrain in north Welsh, the wave drag amounted to 4 dyne/cm², of which 3 dyne/cm² probably acted on the atmosphere above 20 km.

Lilly (1972) reported that “From aircraft traverses through moderate amplitude waves in the Front Range we have commonly obtained direct stress measurements of between 5 and 10 dyne/cm² in the troposphere averaged over horizontal distances of 100-200 km.”

Smith (1978) determined the pressure drag on the Blue-ridge Mountain in the central Appalachians. During the first two weeks of January 1974 he observed several periods with significant wave drag with pressure differences typically of the order of 50 N/m² across the ridge.

Palmer et al. (1986) has pointed out the general westerly bias of the global general circulation models (GCM). They also pointed out that one way to reduce this general westerly bias is to incorporate the gravity wave drag parameterization scheme in the GCM. Gravity wave drag parameterization scheme proposed by Palmer et al. (1986), McFarlane (1987) has reduced the westerly bias mainly in stratosphere.

Recently, lee waves have drawn the new attention from the viewpoint of the atmospheric momentum budget. Iwasaki et al. (1989) introduced a new type of gravity wave drag parameterization scheme to improve the tropospheric westerly bias by including the effects of these tropospheric-trapped lee waves.

Duran (1992), using a simple 2-D model, showed the importance of non-hydrostatic trapped lee waves in the troposphere.

Satomura and Bougeault (1994) used a 2-D, non-hydrostatic, compressible model to simulate the airflow over the Pyrenees in connection with two lee wave events during PYREX experiment. In both cases, the simulated
downward momentum fluxes agree well with the observed fluxes around 4 km height. The simulated fluxes in their study were almost constant with height. The over estimation of simulated momentum flux in the upper half of the atmosphere was suggested to be due to the time evolution of the mean wind and the lateral momentum flux divergence found in the atmosphere.

Vosper and Mobbs (1998) showed that for steady waves, in the absence of dissipation, the vertical fluxes of both the components of horizontal momentum are constant with height.

In India, studies on orographically excited internal gravity wave have done by Das (1964), Sarker (1965), De (1971, 1973), Sarker et al. (1978), Sinha Ray (1988), Kumar et al. (1995). These studies mainly aimed at the computation of wavelength and perturbation streamline displacement, vertical velocity associated with orographically excited internal gravity wave. Dutta (2001) studied momentum/energy flux associated with mountain wave across Mumbai-Pune section of the Western Ghats in an idealized air stream. He showed that both the fluxes were independent of height and the half width of the bell shaped part of the barrier. Further it was shown that the plateau portion of the section does not contribute to the above fluxes.

In a recent study by Wells et al. (2005), the impact of rotation on the orographic drag due to flow across and around a wide mountain have been investigated. In their study, a series of idealized model experiments with a larger range of ridge lengths were performed. Implications of the results obtained for NWP parametrizations of sub-grid-scale orographic drag have been discussed.

Dutta and Naresh (2005) studied momentum and energy flux associated with mountain wave across the Assam-Burma hills. Their study shows that a long valley acts as a source in the atmospheric momentum budget and a sink in the energy budget. In these studies the vertical variation of wind and stability were not considered.

Smith et al. (2006) investigated the behaviour of the resolved Alpine pressure drag with model horizontal resolution using seven mountain wave case-studies from the Mesoscale Alpine Programme (MAP). Three independent modelling systems in their study showed that the magnitude of the resolved surface pressure drag increased monotonically as the horizontal model resolution increased from 125 km down to a few kilometres. They have also shown that a substantial proportion of the drag still need to be parameterized in the next generation of mesoscale models, with horizontal resolutions of about 5 km.

It appears from the foregoing discussions that, although there are some studies on the parameterization of momentum and energy flux associated with orographically excited internal gravity waves with realistic vertical profile of background wind and stability, hardly any such study has been done in India.
Objective of the present study is to propose a dynamical model for parameterizing energy flux and momentum fluxes associated with IGW across a 3-D mesoscale orographic barrier. In Fig. 1, schematically it has been shown how this proposed dynamical model can be used in a large scale/regional scale/meso(α) scale model for a better parameterization of momentum/energy flux in a region near a meso-scale barrier.

2. Data

Radio sonde data on selected dates of Santacruz and Agartala have been used to study near the WG region and the KJ hills respectively, because these two stations are located at far upstream distance with respect to westerly basic flow across the WG and with respect to southerly basic flow across the KJ hills. These data have been collected from the Archival of India Meteorological Department.

3. Model

The dynamical model used in this study is identical to Dutta (2005, 2007) where following assumptions considered:

(i) Steady state flow
(ii) Non-hydrostatic
(iii) Adiabatic basic flow (Dry or moist)
(iv) Laminar
(v) Boussinesq

Following Dutta (2005, 2007), the linearised governing equations, under the above assumptions, are subjected to double Fourier transformation and then from the transformed equations \( \tilde{u}, \tilde{v}, \tilde{p}, \tilde{\theta} \) are eliminated to obtain the following vertical structure equation

\[
\frac{\partial^2 \tilde{w}_l}{\partial z^2} + \left\{ f(k,l,z) - \kappa^2 \right\} \tilde{w}_l = 0
\]

where,

\[\tilde{w}(k,l,z) = \frac{\rho_0(0)}{\rho_0(z)} \tilde{w}_l(k,l,z),\]

\[
f(k,l,z) = \frac{N^2_m k^2}{\sigma^2} - \frac{1}{\sigma} \frac{d^2 \sigma}{dz^2} - \frac{1}{\sigma} \frac{d \sigma}{dz} \frac{dp_0}{dz}
+ \frac{1}{4 \rho_0^2 \left( \frac{dp_0}{dz} \right)^2} - \frac{1}{2 \rho_0} \frac{d^2 \rho_0}{dz^2}.
\]

\[
\kappa^2 = k^2 + l^2
\]

\[
\sigma = U k + V l
\]

\[
(k,l) = \text{Horizontal wave number vector}
\]

\[
N^2_m = \frac{g}{\beta_\theta} \frac{d \bar{\theta}}{dz}
\]

\[
\beta = 1 - \frac{L q_s}{C_p} \left( \frac{e + \gamma q_s}{e + q_s} \right).
\]

Detail derivations for the expressions of \( \beta, N^2_m \) are available in Dutta (2007). The meanings of different symbols have been given in appendix.

The Eqn. (1) is solved using the following boundary conditions:

(i) At the lower boundary, i.e., at the surface \( z = 0 \) airflow is tangential to the ground surface, i.e.,
\[w'(x,y,0) = U(0) \frac{\partial \eta'(x,y,0)}{\partial x} + V(0) \frac{\partial \eta'(x,y,0)}{\partial y},\]
where, \( \eta'(x,y,0) \) is the vertical streamline displacement at \( z = 0 \).

Since at the lower boundary \( z = 0 \) the streamline pattern follows the terrain, hence \( \eta'(x,y,0) = h(x,y) \), where \( z = h(x,y) \) is the three-dimensional profile of the orographic barrier.

So, \( \tilde{w}(k,l,0) = \tilde{w}_l(k,l,0) = i \sigma(k,l,0) \tilde{h}(k,l) \)

In this study the three dimensional profile has been approximated by an ellipsoid, given by

\[
h(x,y) = \frac{H}{1 + \left( \frac{x}{a} \right)^2 + \left( \frac{y}{b} \right)^2}
\]

Here, \( x \)-axis is parallel to the minor ridge axis and \( y \)-axis is parallel to the major ridge axis of the elliptical barrier. For the WG, major axis being north-south oriented, \( x \)-axis points towards east, \( y \)-axis points towards
north and z-axis points vertically upwards. For the KJ hills, major axis being east-west oriented, x-axis points towards north, y-axis points towards west and z-axis points vertically upwards. Thus for the WG, u-component of wind is towards east (westerly), v-component towards north (southerly) and for the KJ hills, u-component of wind is towards north (southerly), v-component towards west (easterly).

With above expression for \( h(x, y) \), we have, \( \tilde{h}(k, l) = 2\pi abHK_0(\sqrt{a^2k^2 + b^2l^2}) \), \( K_0 \) being the Bessel function of second kind (Dutta et al., 2002). Although we have taken an analytical profile for the basic state wind and temperature field, \( \psi(k, l, z) \) is an arbitrary function satisfying equations (1), (2) and (3), and also it’s value above the upper boundary is 1. Following Dutta (2005), \( \psi(k, l, z) \) has been computed numerically at different discrete vertical levels, separated by a distance \( d = 0.25 \) km, for a given wave number vector \( (k, l) \).

Again it can be shown that

\[
\bar{p}(k, l, z) = \frac{i\rho_0(z)}{\kappa^2} \left[ \frac{dz}{\sigma^2} w(k, l, z) - \sigma \frac{\partial \sigma}{\partial z} \right]
\]

and the energy flux vector, \( \vec{E} \), is given by

\[
E_z = p'w' \text{ at any vertical level at any model horizontal grid point can be computed, where, ( ) indicates the average over the surface area 'S' of the barrier. Below it is shown that } S = \frac{\pi^2(a+b)H}{2}.
\]

At any height, \( z \) the elliptical contour is given by \( \frac{x^2}{\alpha^2} + \frac{y^2}{\beta^2} = 1 \), where \( \alpha^2 = a^2 \left( \frac{H}{z} - 1 \right) \) and \( \beta^2 = b^2 \left( \frac{H}{z} - 1 \right) \). So, the area ‘dS’ of a strip of thickness \( dz \) at this height is the product of \( dz \) and the perimeter of

for the WG, we have found that

\[
\tilde{w}(k, l, z) = \sqrt{\frac{\rho_0(0)}{\rho_0(z)}} i\sigma(k, l, 0) \tilde{h}(k, l) \frac{\psi(k, l, z)}{\psi(k, l, 0)} \]

(4)

and hence,

(5)
the above ellipse is given by $dS = \pi (a + \beta)dz$. Hence the total surface area of the barrier is

$$S = \int_{0}^{H} \pi (a + \beta)dz = \pi (a + b)\left(\frac{H}{z} - 1\right)dz = \frac{\pi^2 (a + b)H}{2}.$$

4. **Discussions**

In this section some selected cases for the WG as well as for the KJ hills have been studied. Sarker (1965, 1967) showed that the air stream during winter/pre-monsoon and during the summer monsoon is favourable for the formation of lee waves across the WG. Also De (1973), Sinha Ray (1988) and Sarker et al. (1978) have shown that the air stream during the summer monsoon is
favourable for the formation of lee waves across the KJ hills. Accordingly in this section four cases have been studied, viz., lee wave across the WG during winter/pre-monsoon, lee wave across the WG during summer Monsoon and two lee wave cases across the KJ hills during summer monsoon.

(i) For cases of lee waves across the Western Ghats (WG) during winter/pre-monsoon and summer monsoon.

Case-1: For the lee wave across the WG during winter/pre-monsoon (on 6 March 1965)

The vertical profiles of $U$, $V$ and $T$ are shown in Fig. 2(a), which are based on the 0000 UTC and 1200 UTC Radio sonde data of Santacruz of 6th March 1965.

Using the above profile the vertical energy flux ($E_z$) and momentum fluxes ($\tau_{zx}$ and $\tau_{zy}$) at different levels have been computed.
The vertical profile of energy flux ($E_z$) has been shown in Fig. 2(b). From this figure it can be seen that, whether '$V$' is present or not, above $z = 3.5$ km, $E_z$ is practically invariant in the vertical. It can also be seen that it remains invariant in the vertical at lower levels up to $z = 1$ km in the absence of '$V$' component. Divergence/convergence in $E_z$ can be seen in the layer up to $z = 3.5$ km in presence of '$V$' and in the layer from $z = 1$ km to $z = 3.5$ km, in the absence of '$V$'. It is to note that presence of the vertical shear of the basic flow and the presence of '$V$' makes $E_z$ vertically downward, whereas in the absence of vertical shear of basic flow it is upward.

The vertical profiles of westerly-momentum flux ($\tau_{xw}$) and southerly-momentum flux ($\tau_{ys}$) have been shown in Figs. 2(c&d). From these figures it can be seen that, in the absence of '$V$', they are practically zero and invariant in the vertical. Presence of '$V$' only makes them
divergent/convergent. Vertically upward $\tau_{zx}$ can also be seen in presence of \textquote{\textit{V}}, which is generally vertically downward in case of a constant \textquote{\textit{U}} and $V = 0$.

Case-2: For the lee wave across the WG during Southwest monsoon (on 3 July 2001)

The vertical profiles of $U$, $V$ and $T$, based on average of 0000 UTC and 1200 UTC Radio sonde data of Santacruz of 3rd July 2001, are shown in Fig. 3(a). Since during July the westerly airflow impinging on WG is moist, hence in this case, $T(z)$ has been approximated by the pseudo-adiabatic line through the surface dry bulb temperature. In Dutta (2005), it has been shown that this vertical profile of wind is favourable for giving mountain wave.

The vertical profile of energy flux ($E_z$) in watt/msq is shown in Fig. 3(b). First it can be seen that energy flux is upward. It can also be seen that, whether \textquote{\textit{V} (mean southerly)} is present or not, $E_z$ is divergent in the vertical up to 0.5 km and above this level it is practically non divergent.

The vertical profile of westerly-momentum flux ($\tau_{zx}$) in N/msq is shown in Fig. 3(c). First it can be seen that, this flux is downward whether \textquote{\textit{V}} component is present or not. It can also be seen that magnitude of this flux has enhanced in presence of this component of basic flow. It is observed from this figure that $\tau_{zx}$ is convergent in a shallow layer near the surface above which it is divergent. Convergence/divergence of $\tau_{zx}$ is enhanced in presence of \textquote{\textit{V}}.

The vertical profile of southerly-momentum flux ($\tau_{zy}$) in N/msq is shown in Fig. 3(d). Similar to $\tau_{zx}$, this flux is also downward throughout the vertical domain. We find from this figure that irrespective of the \textquote{\textit{V}} component of basic flow, $\tau_{zy}$ is convergent in the layer from surface to 2 km, above which it is divergent. In the absence of \textquote{\textit{V}}, it is again convergent in a shallow layer near the top of vertical domain.

\textit{(ii)} For cases of lee waves across the Khasi-Jayantia hills (KJ hills) during summer monsoon:

Although in the model computation of the following cases for the KJ hills, $U$ stands for southerly component and $V$ stands for easterly component of basic flow, while preparing the vertical profile of fluxes care has been taken so that, as usual $\tau_{zx}$ represents vertical flux of westerly momentum and $\tau_{zy}$ represents that of southerly momentum.

Case-3: For the lee wave across the KJ hills during Southwest monsoon (on 9 June 2000)

The vertical profiles of $U$, $V$ and $T$ are shown in Fig. 4(a), based on the 0000 UTC and 1200 UTC Radio sonde data of Agartala of 9th June 2000. In this case $V$ is negative at all levels. Since during this season the southerly flow across KJ hills is neutral with respect to moist adiabatic lapse rate, hence following Sarker (1967), De (1971) here also we have approximated $T(z)$ by the pseudo-adiabat through the surface dry bulb temperature.

The vertical profiles of $E_z$, $\tau_{zx}$ and $\tau_{zy}$ are shown in Figs. 4(b-d).

It is seen from Fig. 4(b) that the magnitude of $E_z$ is one order larger than those in cases 1 and 2. It can be seen that in absence of \textquote{\textit{V}} component energy flux is non-divergent and it is almost nil throughout vertical. In presence of this component of basic flow, we find layers with divergent/convergent energy flux. Energy flux is divergent in the layer from surface to 2.5 km, above this level it is convergent up to 3 km and then it is non-divergent in the remaining vertical domain.

From Figs. 4(c&d), in absence of \textquote{\textit{V}} component, we find almost non-divergent vertical flux of both component of horizontal momentum ($\tau_{zx}$ and $\tau_{zy}$) and also they are practically nil throughout the vertical. First of all both the fluxes are downward in the presence of this component. In presence of this component we find different layers with divergent/convergent momentum flux. We find divergent $\tau_{zx}$ in the layer from surface to 1.5 km, above this level it is convergent up to 3 km, then it is non-divergent up to about 3.5 km and then again it is divergent. $\tau_{zy}$ is found to be divergent in the layer from surface to 2.5 km, above this level it is convergent up to 3 km, then it is non-divergent.

Case-4: For the lee wave across the KJ hills during Southwest monsoon (on 1 August 2000)

The vertical profiles of $U$, $V$ and $T$ are shown in Fig. 5(a), based on the 0000UTC and 1200UTC Radio sonde data of Agartala of 1st August 2000. Similar to previous case, in this case also $V$ is negative at all levels. Following similar argument as in the previous case, here also we have approximated $T(z)$ by the pseudo-adiabat through the surface dry bulb temperature.

The vertical profiles of $E_z$, $\tau_{zx}$ and $\tau_{zy}$ are shown in Figs. 5(b-d).
It is seen from Fig. 5(b) that, the presence of ‘$V$’ component has enhanced the magnitude of $E_z$. In presence of this component of the basic flow, the $E_z$ is divergent in the layer between 0 to 1 km and between 2 to 3 km. $E_z$ is non-divergent from 1 to 2 km and from 3 to 4 km. In the absence of this component, $E_z$ is divergent in the layer from $z = 1$ km to 3 km and non-divergent above it.

From Fig. 5(c) it is seen that $\tau_{xz}$ is upward and is nearly non-divergent in the absence of ‘$V$’ component. In presence of this component $\tau_{xz}$ is downward and also there are layers with divergent/convergent $\tau_{xz}$. We find very strong convergence of $\tau_{xz}$ in the lowest layer up to about 0.5 km above which $\tau_{xz}$ is divergent.

Fig. 5(d) shows strong convergence of $\tau_{zy}$ in the lowest layer up to about 0.5 km above which it is divergent.

5. Conclusions

(i) The above model is capable of computing, the vertical energy flux ($E_z$) and the components of eddy
horizontal momentum flux in the direction parallel to the
major ridge (τ_{zx}) and normal to it (τ_{zx}), at all levels.

(ii) Computation shows that all the fluxes vary in the
vertical. The vertical variation is not uniform with height.
In some layers fluxes are upward and somewhere it is
downward. But in most cases energy flux is upward and
momentum fluxes downward. Direction of fluxes depends
on the vertical profiles of the wind and temperature of
basic flow for respective cases.

(iii) However, the influence of the ‘V’ component of
basic wind has been observed in all cases. This component
makes the energy flux or momentum flux divergent/convergent in the vertical in most of the cases.

(iv) For the WG, τ_{zx} is invariant in the vertical in
absence of ‘V’ component. Other fluxes are invariant in
the vertical in some layers. For the WG, whether this
component is present or not, an upper layer is found where
all the vertical fluxes are non-divergent.

(v) From the cases studied it is found that the magnitude
of energy flux across the KJ hills is one order more than
that across the WG.

(vi) For the KJ hills, in absence of ‘V’ component τ_{zx} is
non-divergent, presence of ‘V’ component makes it
divergent/convergent.

Acknowledgements

Author expresses his sincere thanks to DDGM (WF)
and DDGM (Trg.) for their kind support to carry out this
work. Author is grateful to his guide and teacher Dr. U. S.
De, former ADGM (R), IMD for his useful suggestions.
Author is thankful to all the staff members of CTI for their
co-operations.

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**Appendix**

**List of symbols**

<table>
<thead>
<tr>
<th>Symbol</th>
<th>Meaning</th>
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<tbody>
<tr>
<td>$x$</td>
<td>Horizontal co-ordinate in the direction normal to the major ridge of the elliptical barrier.</td>
</tr>
<tr>
<td>$y$</td>
<td>Horizontal co-ordinate in the direction parallel to the major ridge of the elliptical barrier.</td>
</tr>
<tr>
<td>$z$</td>
<td>Vertical co-ordinate pointing towards local vertical.</td>
</tr>
<tr>
<td>$U$</td>
<td>Horizontal component of basic flow along x-axis.</td>
</tr>
<tr>
<td>$V$</td>
<td>Horizontal component of basic flow along y-axis.</td>
</tr>
<tr>
<td>$u'$</td>
<td>Component of perturbation wind along x-axis.</td>
</tr>
<tr>
<td>$v'$</td>
<td>Component of perturbation wind along y-axis.</td>
</tr>
<tr>
<td>$w'$</td>
<td>Component of perturbation wind along z-axis.</td>
</tr>
<tr>
<td>$\rho_0(z)$</td>
<td>Density of basic flow at level ‘$z$’.</td>
</tr>
<tr>
<td>$q_s$</td>
<td>Basic state saturated mixing ratio of water vapour at level ‘$z$’.</td>
</tr>
<tr>
<td>$\tilde{\theta}_e$</td>
<td>Equivalent potential temperature of basic flow at level ‘$z$’.</td>
</tr>
<tr>
<td>$p'$</td>
<td>Perturbation pressure.</td>
</tr>
<tr>
<td>$\rho'$</td>
<td>Perturbation density.</td>
</tr>
<tr>
<td>$\theta_e'$</td>
<td>Perturbation equivalent potential temperature</td>
</tr>
<tr>
<td>$\tilde{u}', \tilde{v}', \tilde{w}', \tilde{\rho}, \tilde{\rho}, \tilde{\theta}$</td>
<td>2-D Fourier transform of $u', v', w', \rho', \rho$ and $\theta'$ respectively.</td>
</tr>
<tr>
<td>$\tilde{T}$</td>
<td>Absolute temperature of the basic flow.</td>
</tr>
<tr>
<td>$k$</td>
<td>Horizontal wave number along x-axis.</td>
</tr>
<tr>
<td>$l$</td>
<td>Horizontal wave number along y-axis.</td>
</tr>
<tr>
<td>$\kappa = \sqrt{k^2 + l^2}$</td>
<td>Magnitude of horizontal wave number vector.</td>
</tr>
<tr>
<td>$C_p, C_v$</td>
<td>Specific heats of gas at constant pressure and at constant volume respectively.</td>
</tr>
<tr>
<td>$\gamma$</td>
<td>$C_p/C_v$</td>
</tr>
<tr>
<td>$z = h(x, y)$</td>
<td>Analytical expression for 3-D elliptical barrier.</td>
</tr>
<tr>
<td>$\tilde{h}(k,l)$</td>
<td>2-D Fourier transform of $h(x, y)$.</td>
</tr>
<tr>
<td>$\psi(k,l,z)$</td>
<td>Arbitrarily defined function, satisfying the same upper boundary condition as $\tilde{w}_1(k,l,z)$</td>
</tr>
<tr>
<td>$\sigma(k,l,z) = kU(z) + lV(z)$</td>
<td>Component of the basic flow along the wave number vector $(k,l)$.</td>
</tr>
</tbody>
</table>